

A magnitude 7.0 earthquake occurred 181.3 km (112.6 miles) SE of Davao, and 83.8 km (52.1 miles) SE of the city of Pondaguitan at a depth of 60.1 km (37.3 miles).

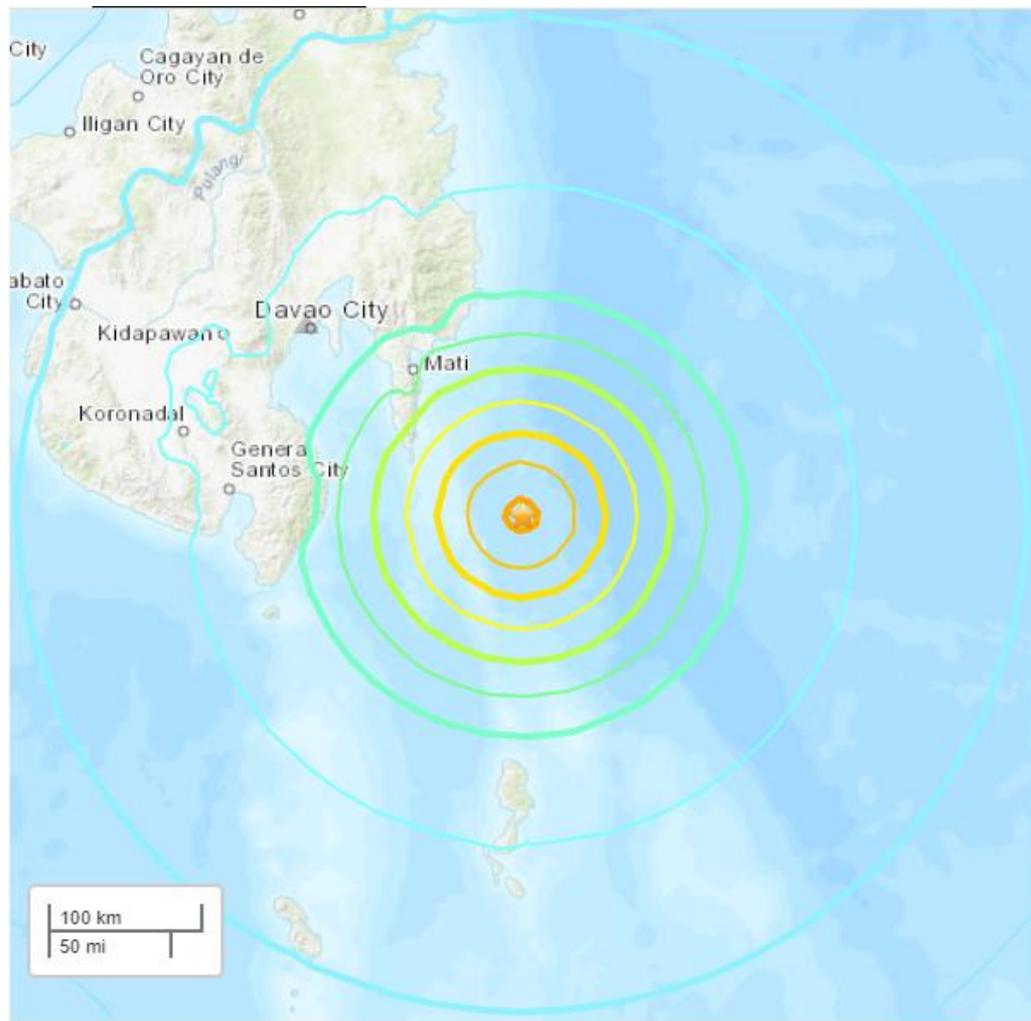
There are no immediate reports of casualties or damage.



The Modified-Mercalli Intensity scale is a twelve-stage scale, from I to XII, that indicates the severity of ground shaking.

The area nearest the earthquake experienced strong shaking from this earthquake.

Modified Mercalli Intensity	Perceived Shaking
X	Extreme
IX	Violent
VIII	Severe
VII	Very Strong
VI	Strong
V	Moderate
IV	Light
II-III	Weak
I	Not Felt



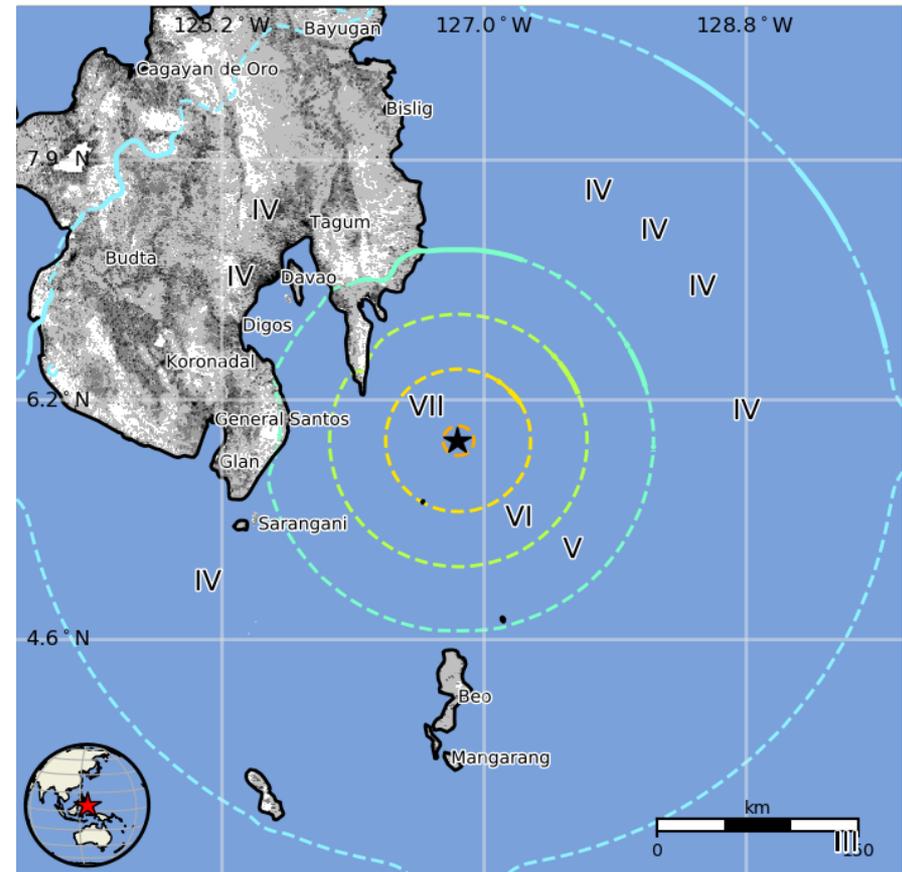
USGS Estimated shaking Intensity from M 7.0 Earthquake

The USGS PAGER map shows the population exposed to different Modified Mercalli Intensity (MMI) levels.

The USGS estimates that 16,000 people felt strong shaking from this earthquake.

I	Not Felt	0 k*
II-III	Weak	4,024 k*
IV	Light	13,936 k
V	Moderate	539 k
VI	Strong	16 k
VII	Very Strong	0 k
VIII	Severe	0 k
IX	Violent	0 k
X	Extreme	0 k

Population Exposed to Earthquake Shaking



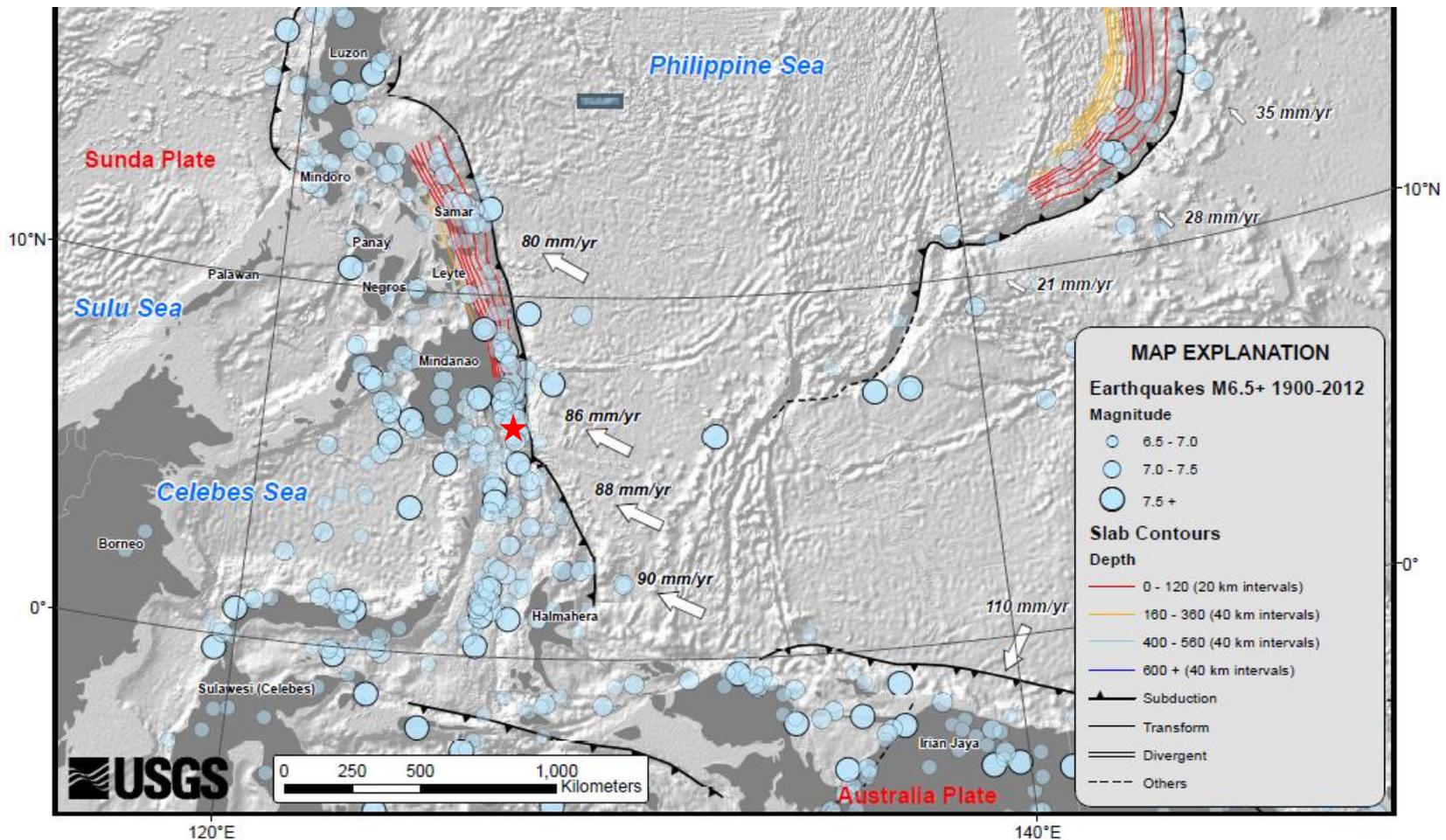
The color coded contour lines outline regions of MMI intensity. The total population exposure to a given MMI value is obtained by summing the population between the contour lines. The estimated population exposure to each MMI Intensity is shown in the table.

Image courtesy of the US Geological Survey

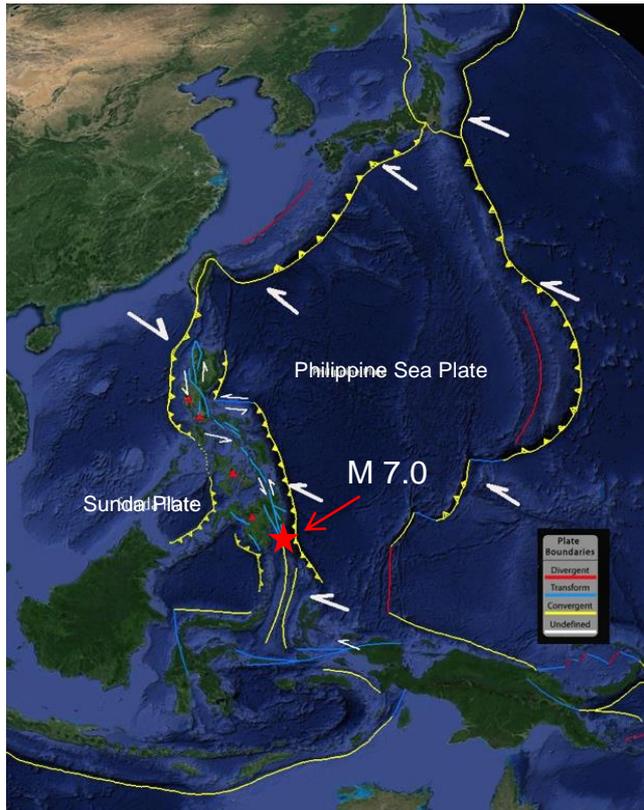
Magnitude 7.0 PHILIPPINES

Saturday, December 29, 2018 at 03:39:09 UTC

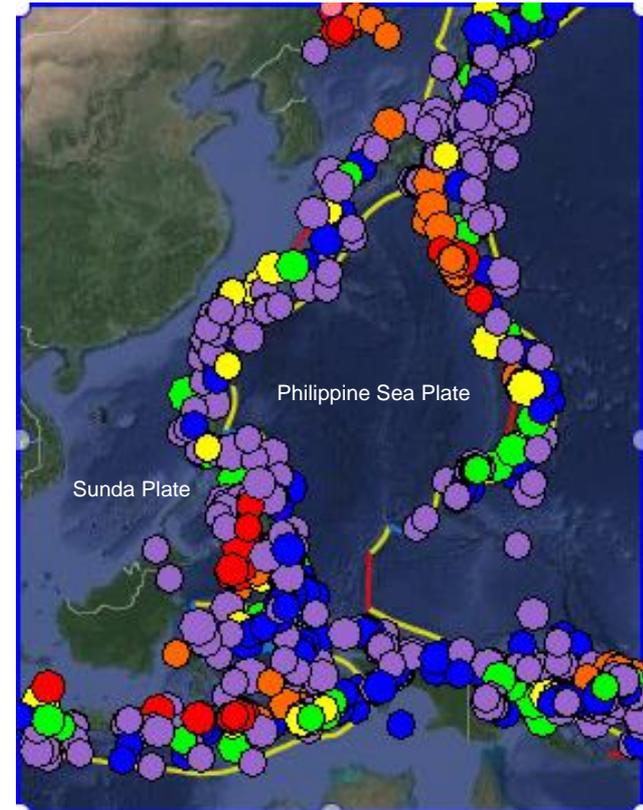
At the latitude of this earthquake, the Philippine Sea Plate moves towards the west-northwest with respect to the Sunda Plate at a rate of approximately 8.6 cm/yr. The Philippine Sea Plate subducts beneath the Philippine Islands at the Philippine Trench at the location of this earthquake.



Along its western margin, the Philippine Sea Plate is complicated where it converges with, and dives beneath the Sunda Plate. Caught in the crunch, the Philippines archipelago has oceanic plates subducting beneath both its east and west sides, and the arc complex itself is marked by active volcanism (red triangles), as well as high seismic activity.



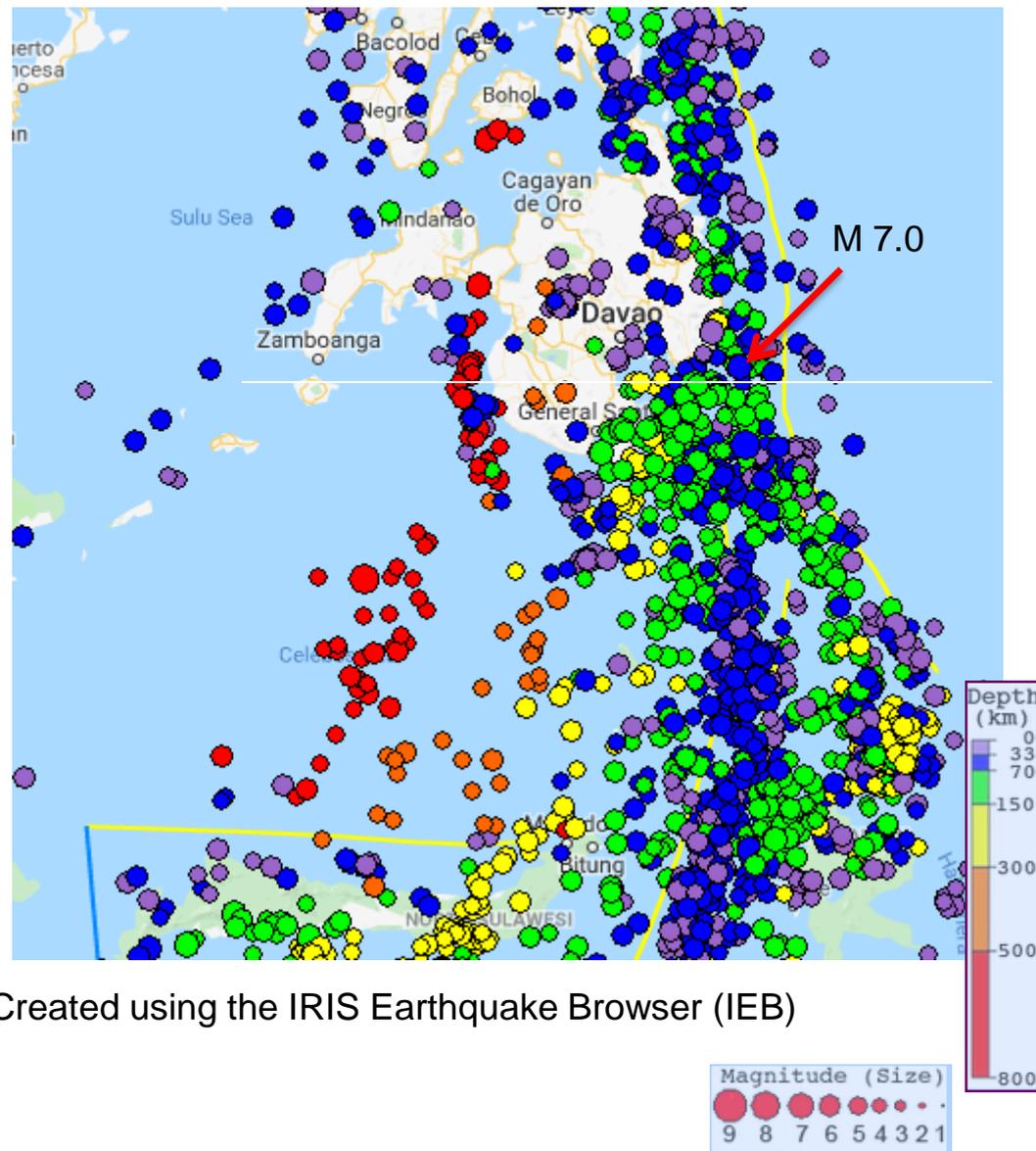
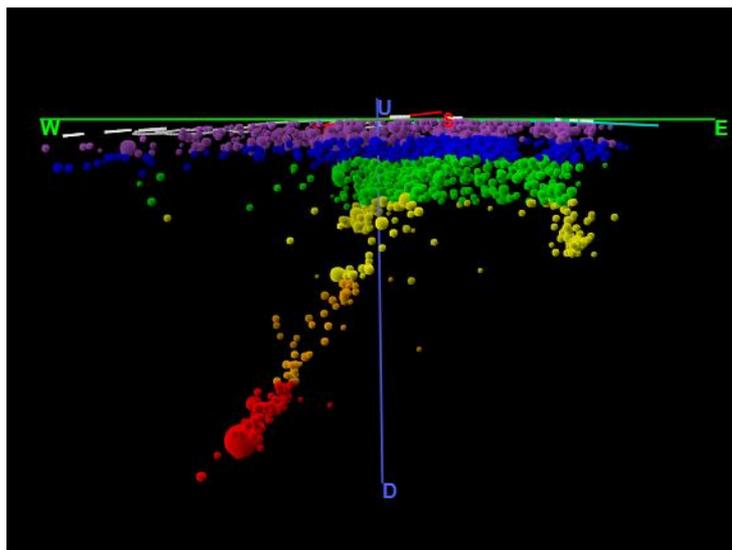
Simplified tectonic contacts



Magnitude 6-8 earthquakes 2000-2018

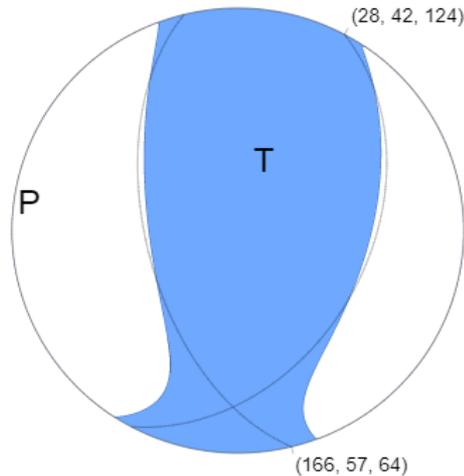
This map shows historical seismicity in this region. Earthquakes are color-coded by depth as shown in the legend in the lower right corner. Depths of earthquakes increase from east to west across the subduction zone boundary.

A 3D cross section through the earthquake is shown below.



Created using the IRIS Earthquake Browser (IEB)

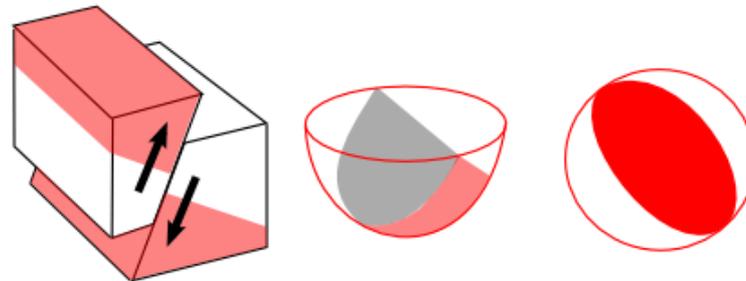
The focal mechanism is how seismologists plot the 3-D stress orientations of an earthquake. Because an earthquake occurs as slip on a fault, it generates primary (P) waves in quadrants where the first pulse is compressional (shaded) and quadrants where the first pulse is extensional (white). The orientation of these quadrants determined from recorded seismic waves determines the type of fault that produced the earthquake.



USGS W-phase Moment Tensor Solution

The tension axis (T) reflects the minimum compressive stress direction. The pressure axis (P) reflects the maximum compressive stress direction.

Reverse/Thrust/Compression



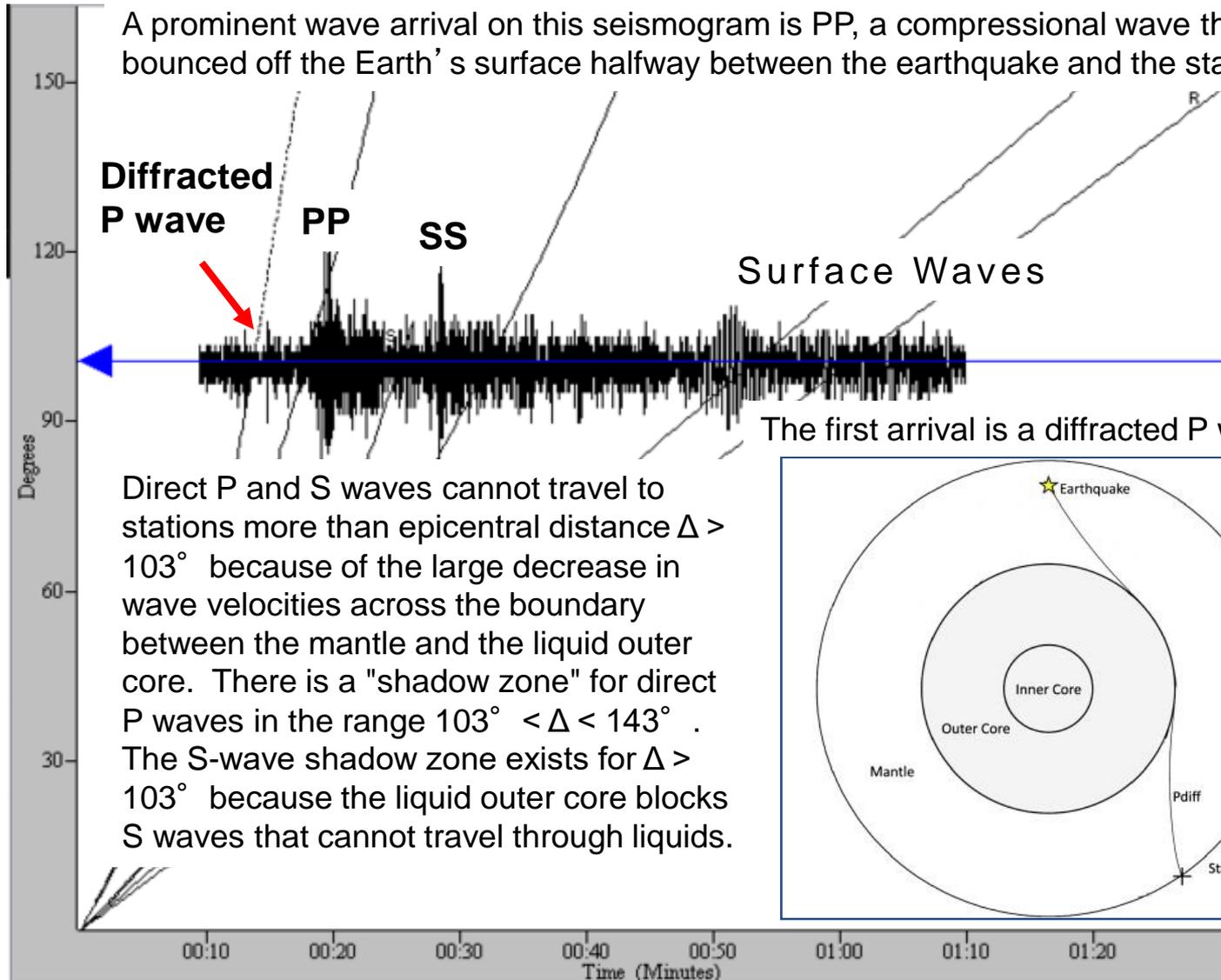
In this case, the focal mechanism indicates this earthquake occurred as the result of thrust faulting.

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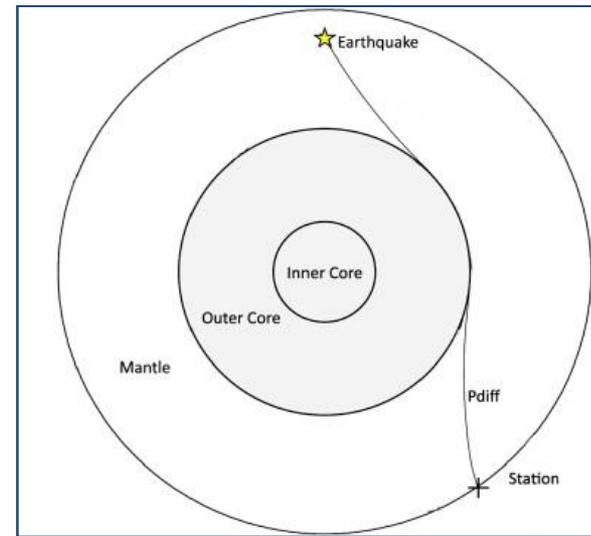
The record of the earthquake in Bend, Oregon (BNOR) is illustrated below.
Bend is 11,209 km (6964 miles, 100.9°) from the location of this earthquake.

A prominent wave arrival on this seismogram is PP, a compressional wave that bounced off the Earth's surface halfway between the earthquake and the station.



The first arrival is a diffracted P wave.

Direct P and S waves cannot travel to stations more than epicentral distance $\Delta > 103^\circ$ because of the large decrease in wave velocities across the boundary between the mantle and the liquid outer core. There is a "shadow zone" for direct P waves in the range $103^\circ < \Delta < 143^\circ$. The S-wave shadow zone exists for $\Delta > 103^\circ$ because the liquid outer core blocks S waves that cannot travel through liquids.



Animation explaining the seismic shadow zone.

Epicentral distance is the angle formed by the intersection of the line from the earthquake to Earth's center with the line from the observing point to the Earth's center.

S waves are observed up to a distance of 104° from an earthquake, but direct S waves are not recorded beyond this distance.

P waves also have a shadow zone between 104° and 143° .



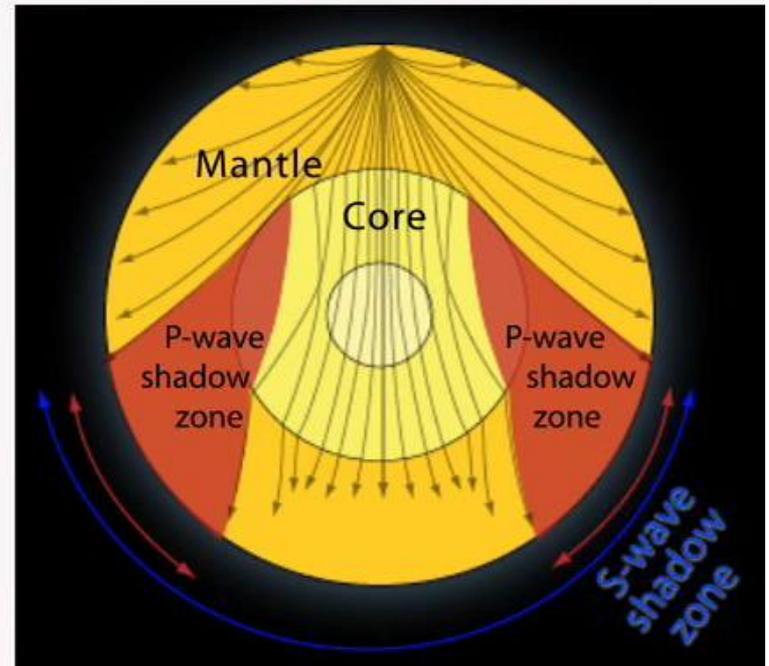
Seismic Shadow Zones



How the mantle and core were determined using the arrival times of direct P and S body waves

P waves (primary) are compressive waves that travel through solids & liquids.

S waves (secondary) are shear waves that travel through solids only.



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